# Global water resources assessment under climatic change in 2050 using TRIP

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**Abstract** Annual water availability was derived from annual runoff estimated by land surface models using Total Runoff Integrating Pathways (TRIP) with  $0.5^{\circ} \times 0.5^{\circ}$  longitude/latitude resolution globally. Global distribution of water withdrawal for each sector in the same horizontal spatial resolution was estimated based on country-base statistics. Based on the framework, future projections of the global water resources considering population growth, climatic change, and the increase of water consumption per capita were carried out. Change in annual runoff was estimated based on the climatic simulation by a general circulation model coupled with TRIP. With the increase of population only, future population under strong water scarcity, when water scarcity index is larger than 0.4, will increase by 90% in 2050 compared to the current situation in 1995. Consideration of the climatic change due to global warming will relax this situation, and only 74% will be under strong water scarcity according to the future projection used in this study.**[key words?]** 

# **INTRODUCTION**

The water resources situation is seriously under pressure, even at present in some regions of the earth, and it is anticipated that the situation will be worse during the first half of the twenty-first century (Cosgrove & Rijsberman, 2000). The estimation of the current level of water stress is important for reliable projections of the severity of the water crisis in the future. However, most of the previous global analysis on water scarcity have been carried out on a country basis or a river basin basis with the exceptions of Takahashi *et al.* (2000), Vörösmarty *et al.* (2000), and Oki *et al.* (2001), who presented water scarcities on a each  $0.5^{\circ} \times 0.5^{\circ}$  longitude and latitude grid for the globe. Considering the importance of global water scarcity, future projections should be evaluated by multiple procedures, models, and methods at multiple organizations, since the reliability of the estimates will be supported if similar results are obtained from different scientific approaches, information, and data processing. As the future projection of global warming has been done under IPCC (Intergovernmental Panel on Climate Change), such an approach to evaluate the future prediction by multiple models/procedures is expected to reduce the uncertainty naturally embodied in the projections.

In this study, water balances estimated by land surface models (LSMs) were used for the global estimates of water availability. Country-based statistics, such as municipal water supply and industrial water use, were assumed to be proportional to the population distribution, and agricultural intake was assumed to be proportional to the irrigated land area. Detailed procedures are described in Oki *et al.* (2001), and future projections of world water resources assessment were carried out using the framework, considering the change in climatic conditions estimated by a general circulation model (GCM).

From these estimates of global annual water supply and demand in 2050, the global distribution of water stress was estimated for every  $0.5^{\circ} \times 0.5^{\circ}$  longitude/latitude grid boxes globally, except for grid boxes which are more than half covered by ice or water. The result was examined with water stress levels in terms of increase or decrease compared to the current situation in 1995, and the dominant factors on the increase were evaluated among population growth, climate change, and the increase of water usage per capita.

# WATER AVAILABILITY IN 2050

There are several methods to be adopted to estimate the water availability in the future. Firstly, the projected water cycles by climate model may be directly used for water resources assessment. This is straightforward, however, model outputs often suffer bias errors, particularly river runoff estimates, and correcting the errors is generally required for quantitative estimation of available water resources in the future. The coarse spatial resolution of the ordinary climate model is repeatedly claimed as for the insufficiency of utilizing model outputs, but the spatial resolution is becoming rapidly higher. Regional climate models may be nested in the global projection of the future climate in order to obtain the future projection with higher spatial resolution. However, the model bias should be removed.

There are ways to downscale the frequency of weather pattern to the climate on regional scales. It assumes the current relationship between weather pattern and climatic variables, such as pressure distribution and precipitation at each point, persists in the future. There is no guarantee to assume the persistence and another uncertainty of the relationship to convert weather pattern to climatic variables is involved in the future projection, in addition to the uncertainty originated from the climate models.

Another way to estimate the future situation under global warming is to use historic climatic data when it was warmer, as a proxy of the future situation. This procedure is also claimed to be ambiguous since, for example, there is no guarantee that the hydrologic cycles are similar for the past warm period and future climate under global warming. It is practically difficult to apply this proxy method for global water resources assessment because it is not easy to obtain global information on water availability in the past.

In this study, the current global distribution of available water resources was estimated, adjusted using GCM simulation on the future climate, and used for global water resources assessment.

For the estimation of the current runoff distribution on a global scale, offline runs by 10 different LSMs were used. Under the Global Soil Wetness Project (GSWP; IGPO (1995)), the global water balance was estimated by 11 LSMs using forcing data, such as precipitation, downward short wave and long wave radiation, wind speed, temperature, humidity, and surface pressure, from ISLSCP (International Satellite Land Surface Climatology Project) (Meeson *et al.*, 1995). TRIP (Total Runoff Integrating Pathways) (Oki & Sud, 1998) was used for the river routing calculations to convert runoff from LSMs into river discharge. Details are described in Oki *et al.* (2001), but basically the estimated annual discharge corresponded fairly well with observation where forcing data to run LSMs were made with enough real observations, such as rainfall (Oki *et al.*, 1999). As seem in Oki *et al.* (2001), the river discharge here is smaller compared to the previous estimates by approximately 20%, and this shortcoming is one of the critical points of the current study for further improvement.

Two methods are applied here to adjust the current river runoff into future river runoff. One method is to use the difference in GCM simulation for current and future climatic conditions, and the other one is to use the ratio. Both methods may suffer technical problems when adjusting current river runoff to the future. The difference method estimates future river runoff  $r_{fd}$  by modifying current river runoff  $r_n$  with the difference of current  $R_n$  and future  $R_f$  runoff simulations by GCM:

$$r_{fd} = r_n + (R_f - R_n) \tag{1}$$

As equation 1 can be written as:

$$r_{fd} = R_f - (R_n - r_n) = R_f - \Delta R \tag{2}$$

with the model bias  $\Delta R = R_n - r_n$ , it implicitly assumes the water balance estimated by the GCM may have a bias error but that is constant through the future. The technical problem is that the negative runoff can be estimated when the model bias is larger than GCM estimates of future river runoff  $\Delta R > R_f$ . Practically negative runoff should be filled with zero runoff.

The ratio method estimates future runoff  $r_{fr}$  by:

$$r_{fr} = r_n \times \frac{R_f}{R_n} \tag{3}$$

and it assumes the systematic error by GCM simulation keeps the ratio constantly.  $r_{fr}$ 

may be estimated tremendously big if current GCM runoff  $(R_n)$  is close to zero, but some future runoff  $(R_f)$  is predicted  $(R_f/R_n$  is huge), and real runoff  $(r_n)$  is not zero. Such a situation occurs mainly in desert regions where the availability of water resources can be neglected; therefore such an unrealistic estimate may be neglected for use in water resources assessment. In this study, both  $r_{fd}$  and  $r_{fr}$  are estimated and used for future water resources assessments.

An atmospheric general circulation model (AGCM) of the Centre for Climate System Study, University of Tokyo and the National Institute of Environmental Studies, Japan (Numaguti *et al.*, 1997) was run with boundary conditions of sea surface temperature in both 1990 and 2060, that are obtained from coupled oceanatmospheric general circulation model (AOGCM) run for increasing carbon dioxide conditions. The horizontal resolution of the model simulation was carried out with T106, which corresponds to approximately 100-km grid boxes globally. Daily runoff estimates from AGCM outputs were taken and routed using TRIP. The annual mean discharges in 1990 and 2060 were regarded as current ( $R_n$ ) and future ( $R_f$ ) river discharge. The difference  $R_f - R_n$  and ratio  $R_f/R_n$  are shown in Figures 1 and 2.



**Fig. 1** The difference of the simulated river discharge for 1990  $(R_f)$  and 2060  $(R_{nf})$ .



From this particular simulation, the Asian monsoon is enhanced due to the enhancement of the temperature difference between the Indian Ocean and the Eurasian continent. River discharge is expected to increase in the Indian sub-continent through the western part of the Indo-China Peninsula. Northern China, where water is relatively scarce at present, is also expected to increase its potentially available water resources, and a decrease is projected in between. Most of Europe is projected to decrease its annual river discharge. In contrast, the model simulation projects an increase in Central to Western Africa. The ratio  $R_f/R_n$  is sometimes extremely large near the dry region in Fig. 2, and the small increase of river runoff is exaggerated, as expected. The future river runoff estimated using difference  $r_{fd}$  is negative in some parts of the western US (not shown). The difference of these two kinds of future estimates  $(r_{fr} - r_{fd})$  is shown in Fig. 3. Since the area of 0.5° grid cell is approximately 2500 km<sup>2</sup> in tropical regions, 5000 × 10<sup>6</sup> m<sup>3</sup> per 0.5 grid cell corresponds to a difference of 2000 mm year<sup>-1</sup>. Therefore for some regions future projection is highly dependent on the assumption employed for bias error correction of GCM.



Fig. 3 The difference of river discharge in 2050 estimated by difference method  $r_{fd}$  and ratio method  $r_{fr}$ .



Fig. 4 The change in the total water withdrawal per year under the scenario 3 in 2050.

Both estimates of  $r_{fr}$  and  $r_{fd}$  were used to assess the water resources in 2050, together with the future projection of water withdrawals described in the next section.

## FUTURE WATER WITHDRAWAL

The current water withdrawal estimated in Oki *et al.* (2001) was used for the baseline for the future prediction of water withdrawal in the world. Three scenarios were assessed in this study. The first scenario assumes that only the population will increase according to the medium prediction of the United Nations. The river discharge will change due to the climatic variation in the second scenario, in addition to the increase of population. The future river discharges of  $r_{fd}$  and  $r_{fr}$  presented in the previous section were used. The unit withdrawal of municipal and industrial water will increase or decrease because of the economic growth and technology development in the third scenario, in addition to the population growth and climatic change.

The municipal water withdrawal increases, even in scenarios 1 and 2, according to the population growth. In scenario 3, the unit municipal water withdrawal will change with the ratio projected by Takahashi *et al.* (2000). Water use efficiency will increase in North America where unit withdrawal will decrease by approximately 20%. Water use efficiency will be neutral for OECD and Western Europe, but gradually increase in developing countries.

The industrial water withdrawal will change only under scenario 3 with the ratio by Takahashi *et al.* (2000). In scenario 3, it was also assumed that the unit water withdrawal would increase proportional to the growth of GDP predicted by Raskin *et al.* (1995) in each region.

The future projection of irrigation withdrawal in this study is not sophisticated and will need further refinements. From the past 40 years of records it was found that the increase in irrigation area in the world is approximately proportional to the increase in the world population, and it was assumed the situation would continue for the next 50 years. Since the unit irrigation withdrawal was assumed to be constant in each country, as a result, the increase in irrigation withdrawal is proportional to the population growth here.

Figure 4 illustrates the change in the predicted total water withdrawal per year under scenario 3. Basically it reflects the change in population and therefore decreases in developed countries and increases in developing countries.

#### **GLOBAL WATER RESOURCES ASSESSMENT IN THE FUTURE**

The assessment of the water scarcity in this study adopted the index based on the ratio of the annual water withdrawal W to the available annual water Q, and used by UN *et al.* (1997) and Vörösmarty *et al.* (2000). Following Heap *et al.* (1998), the desalinated water resource S was subtracted from W, and the water scarcity index  $R_{ws}$  Falkenmark *et al.* (1989) was derived as:

$$R_{ws} = \frac{W - S}{Q} \tag{4}$$

The global distribution of estimated  $R_{ws}$  is shown in Fig. 5 for scenario 3 using  $r_{fd}$ . Generally the severity of water scarcity is judged as:

 $R_{ws} \leq 0.1$  no water stress;

 $0.1 < R_{ws} \le 0.2$  low water stress;

 $0.2 < R_{ws} \le 0.4$  moderate water stress;

 $0.4 \le R_{ws}$  high water stress

Following these criteria, it is evident from Fig. 5 that water scarcity is severe in the river basins of the Yellow, Indus, Ganges, and Amu-Darya rivers, and in the mid-west of the United States. These distributions are generally similar to the current situation.



Fig. 5 The future projection of the water withdrawal to availability ratio in 2050.



However, looking at the change in ratio of  $R_{ws}$  in the future to the current  $R_{ws}$  (Fig. 6), it is apparent that a drastic change is anticipated in Africa. In the continent, current

Stress level	Current (1995)	Population growth only	+ Climate change		+ Unit demand increase	
			$r_{fd}$	r <sub>fr</sub>	$r_{fd}$	r <sub>fr</sub>
High	1.67	2.66	2.43	2.57	2.51	2.65
Moderate	0.58	0.91	0.93	0.93	0.97	0.95
Low	0.62	0.94	0.96	0.94	1.00	0.96
No	2.78	4.31	4.45	4.28	4.28	4.15

**Table 1** World Population (billion people) under each Water Stress Level projected in 2050. Future river discharge under climatic condition was estimated in two ways  $r_{fd}$  using the difference and  $r_{fr}$  using the ratio of current and future GCM simulations.

water withdrawal is comparatively little, however, rapid increase of water withdrawal is projected according to the population growth. The situation should be more serious than those regions where  $R_{ws}$  will be as high as these African regions, but  $R_{ws}$  is already high because social infrastructure to secure water resources and better water resources management system should already be in place in such regions. However, a sudden increase in  $R_{ws}$  should demand the development of such an infrastructure and better management systems. The situation will become serious if this cannot be achieved.

Impacts of climatic change and economic growth associated with technology development are assessed in Table 1 for  $r_{fd}$  and  $r_{fr}$ . The estimated values for the current situation are also presented in Table 1 for comparison. Basically, the population growth has the largest impact on the increase of people under high water stress level from 1.7 billion now to 2.5, through 2.7 billion in 2050. Since global river runoff increases in the GCM simulation used in this study, the water stress level is somewhat relaxed by climate change effect. This is the opposite result to that by Vörösmarty *et al.* (2000) in which the population under high water stress level increases if the impact of climate change is considered. However, the impact of climate change effect is not as dominant as population growth in their estimates. Of course, in developed countries, population will decrease and the water stress level will be reduced. Therefore the precise prediction of the impact of climate changes should have more relevance in developed countries.

Even though the difference of  $r_{fd}$  and  $r_{fr}$  are large in some regions, the difference in Table 1 is comparatively small. The impact of unit demand increase is also small. These results may suggest that such a presentation showing population under each water stress level is useful to know the relative impact of the factors changing the water supply and demand. However, further indexing and detailed presentation are also necessary to assess the change in water resources situation in the future.

### SUMMARY AND FUTURE RESEARCH NEEDS

A framework to assess the future water resources condition by grid base globally was shown and preliminary results are presented. However, each component of the future projection should be sophisticated, particularly about the irrigation withdrawal. It should be coupled with the changes of land use and climate, such as temperature and precipitation. Urbanization was not explicitly considered in this research; however, it may have a big impact, particularly in developing countries, and may need to be considered separately.

The impact of climatic change may be evaluated more if seasonal change of potential water resources is used for assessment, since global warming will change the snow accumulation and the timing of snow melt. In such time scales, river runoff routing calculations should be incorporated with anthropogenic activities on hydrologic cycles, such as reservoir operations and water withdrawals. According to the increase of human impact on nature, "natural" hydrologic cycles excluding anthropogenic activities are different from "real" hydrologic cycles now. It is important for water resources assessment to deal with "real" hydrologic cycles.

Much more effort should be made to investigate the robust information we can extract from GCM simulations. The ultimate goal should be to utilize the realistic and unbiased simulation of hydrologic cycles by GCM directly for water resources assessment. However, it may not be achieved within several years and we need some proxy method for practical needs, at least for a while.

As for the evaluation of water stress level, it may be necessary to consider the virtual water trade, how much water resources are saved by importing agricultural and cattle products. Including such aspects, an integrated approach to improve the accuracy of future projections on both the natural and social sides of the water resources should be promoted for developing action plans to mitigate the water crisis anticipated in the early stage of the twenty-first century.

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